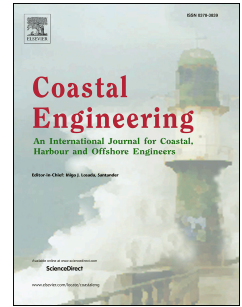


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Combined wave-current induced excess pore-pressure in a sandy seabed: Flume observations and comparisons with theoretical models

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1 **Combined wave-current induced excess pore-pressure in a sandy seabed: Flume**
2 **observations and comparisons with theoretical models**

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8 **ABSTRACT**

9 Waves are coexisting with currents in coastal zones; nevertheless, previous experimental
10 studies for excess pore-pressure responses in a porous seabed were predominantly limited to
11 the wave-only condition. In this study, the combined wave-current induced excess pore-
12 pressures in a sandy seabed were experimentally simulated with a specially-designed flume,
13 which can concurrently generate periodic waves and a following/opposing co-directional
14 current. The effect of a current on the wave profile is firstly examined. The wave steepness is
15 decreased by a following current, but enhanced by an opposing current. Flume observations
16 indicate that, under combined wave-current loading, the wave-induced pore-pressure is
17 increased for the following-current case, but reduced for the opposing-current case. Such
18 wave-current combination effect becomes more significant for shorter wave periods. The
19 variation trend of the excess pore-pressure distribution in the present flume observations is
20 consistent with that of the existing analytical solutions. Nevertheless, due to the existence of
21 wave and/or current boundary layer and non-linearity of wave-current interactions as
22 indicated by the flume observations, certain deviations exist between the flume results for
23 excess pore-pressure and the analytical solutions, which can not be ignored especially for the
24 opposing-current case. The effects of the boundary layer on the combined wave-current
25 induced pore-pressures in the seabed are further highlighted by supplementary numerical
26 simulations. A favorable prediction by the analytical solution would be expected for
27 following-current cases and smaller pore-pressure amplitudes would be obtained for
28 opposing-current cases.

29 **Keywords:** Excess pore-pressure; flume experiment; sandy seabed; combined waves and
30 current; boundary layer; wave-current interactions

31

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32 1. Introduction

33 The evaluation of the wave-induced soil response in marine sediments is particularly
34 important for the design of foundations of offshore installations such as wind turbine
35 foundations (Cuéllar, 2012; Lin et al., 2017), platforms (Bea et al., 1983; Zhang et al., 2017),
36 pipelines (de Groot and Meijers, 1992; Zhou et al., 2015) and breakwaters (Oumeraci, 1994;
37 Zhang and Ge, 1996; de Groot et al., 2006; Liao et al., 2018a, 2018b). Therefore, it is
38 necessary to have a better understanding of the wave-induced pore-pressure in marine
39 sediments.

40 In the past a few decades, numerous analytical solutions have been obtained and several
41 experimental works have been done for wave-induced oscillatory pore-pressure responses.
42 Based on Biot's poro-elastic theory, a few porous models for wave-seabed interactions were
43 ever established under various assumptions (see Sumer, 2014; Jeng, 2018). Among these, the
44 analytical solution by Yamamoto et al. (1978) took into account of compressible pore-water
45 in a compressible isotropic porous seabed with infinite thickness. Madsen (1978) presented a
46 general analytical solution for pore-pressures and effective stresses in a hydraulically
47 anisotropy porous seabed with infinite thickness. With the same framework, Hsu and Jeng
48 (1994) later derived the analytical solution to Biot's equations for the case of finite soil
49 thickness, which can converge to the above solution by Yamamoto et al. (1978) and Madsen
50 (1978) if the soil thickness approaches infinity. The validity of these analytical solutions have
51 been confirmed by both one-dimensional tests using cylindrical-shaped apparatuses
52 (Chowdhury et al., 2006; Liu et al., 2015) and flume experiments (Tsui and Helfrich, 1983;
53 Chang et al., 2007; Zhang et al., 2016; Zhai et al., 2018). A detailed review of the previous
54 investigations on the wave-seabed interaction can be found in Jeng (2003; 2013; 2018).

55 In natural ocean environments, waves are coexisting with currents. The pore-pressure
56 responses of the seabed could be significantly different when a current is considered. To the
57 author's knowledge, Ye and Jeng (2012) were the first ones to study the soil response for the
58 scenario of combined waves and currents. Numerical simulations were conducted based on

59 Biot's poro-elastic dynamic theory ($u-p$ approximation). Their results showed that the
60 maximum relative difference of the pore-pressure between the cases with currents and
61 without currents can reach up to 25%. Zhang et al. (2013) proposed an analytical
62 approximation for the evaluation of the pore-pressure in the seabed under combined waves
63 and currents by adopting an updated wave-induced pressure at the seabed surface. It indicated
64 that the influence of a current on the pore-pressure responses is significant. The full dynamic
65 soil behavior was considered by Liao et al. (2013) and an analytical solution of the pore-
66 pressure responses was derived for an infinite seabed. The parametric study showed that the
67 current with third-order wave loading and full dynamic soil behavior cannot be ignored in the
68 estimation of the wave-induced seabed responses for nearly-saturated soil, long-wave periods,
69 and shallow water. Wen et al. (2016) established a three-dimensional numerical model for
70 pore-pressure response under combined short-crested waves and currents. The numerical
71 results indicated that superimposing a following-current will result in larger pore-pressure in
72 the seabed. Therefore, ignoring a following-current would underestimate the wave-induced
73 seabed instability.

74 As aforementioned analytical and numerical studies indicated, while considering the
75 combined wave-current induced pore-pressure responses in a seabed, the Biot's poroelastic
76 theory (Biot, 1941, 1960) is accepted as the principle of compressible pore fluid flow in a
77 compressive porous medium. The governing equations of seabed responses are the same for
78 wave-only condition and combined wave-current condition. Consequently, the essential
79 difference of the pore-pressures between wave-only condition and combined wave-current
80 condition is induced by the different boundary conditions of pressure at the seabed surface.
81 This highlights the significance of the effect of wave-current interaction on the pressure
82 distributions at the seabed surface. However, despite a substantial amount of knowledge has
83 accumulated about the effect of wave-current interactions on the velocity profiles and
84 turbulence characteristics (e.g. Kemp and Simons, 1982, 1983; Zhang et al., 2014; Tambroni
85 et al., 2015; Singh and Debnath, 2016), little attention has been paid to the effect of wave-
86 current interaction on the pressure distributions at the seabed surface. Moreover, the existing

87 studies with respect to combined wave-current induced pore-pressure responses in the seabed
88 were predominantly limited to deriving analytical solutions and conducting numerical
89 simulations. A systematic experimental study on the excess pore-pressure responses under
90 combined waves and current has not been available in the literature. Note that the “excess
91 pore-pressure” herein denotes the wave-induced pore-pressure relative to the still hydrostatic
92 pressure in the seabed (refer to Yamamoto et al., 1978; Zen and Yamazaki, 1990).

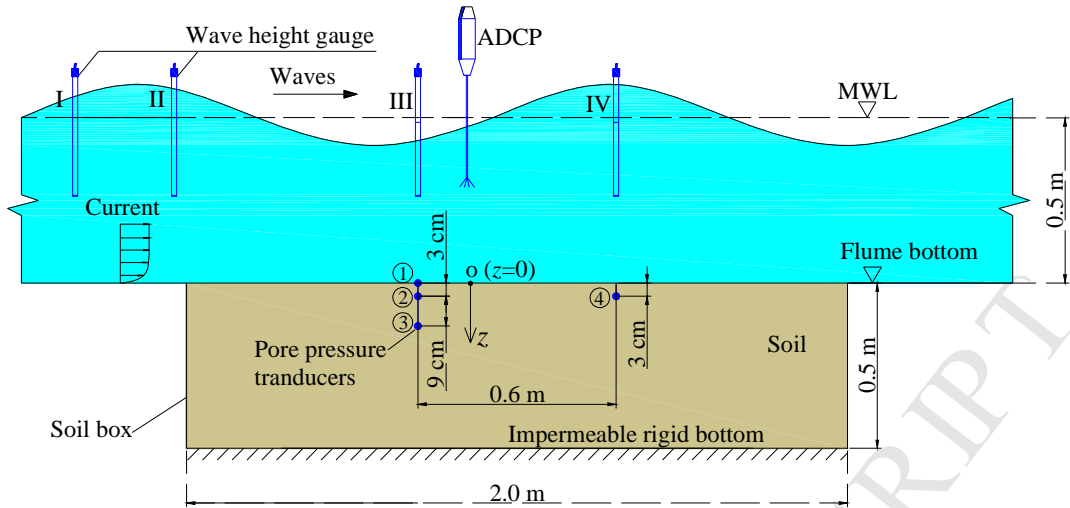
93 In the present study, a series of large flume tests were conducted to investigate the
94 excess pore-pressure responses in a sandy seabed under combined waves and current. To
95 examine the effect of wave-current combination on the excess pore-pressure responses,
96 various magnitudes of the following-current and the opposing-current were superimposed on
97 the waves. The variation of the excess pore-pressure responses with wave period were
98 investigated for the conditions of wave-only and combined waves and current. Moreover, the
99 applicability of the existing analytical solution was examined by comparing the excess pore-
100 pressure distributions between the experimental results and the existing analytical solutions.
101 Several numerical simulations were also carried out to elucidate the significant effects of
102 boundary layer on the combined wave-current induced pore-pressures in the seabed.

103

104 **2. Experimental study**

105 **2.1. Experimental set-up**

106 The combined wave-current induced excess pore-pressure responses in a sandy seabed
107 were experimentally simulated with a specially-designed flume, which can concurrently
108 generate periodic waves and a following/opposing co-directional current. The major frame of
109 the flume is 52.0 m in length, 1.0 m in width and 1.5 m in depth, in the middle of which a
110 soil-box of 2.0 m (length)×0.5 m (depth)×1.0 m (width) was constructed for the sand-bed
111 preparation, as illustrated in Fig. 1.



112

113 Fig. 1. Schematic diagram of the flume tests for combined wave-current induced excess pore-pressure
 114 in a sandy seabed.

115

116 A sandy seabed was prepared with a sand-raining device, whose main physical
 117 properties are listed in Table 1. Four miniature pore-pressure transducers (PPTs) with model
 118 number of GE Druck PDCR 81 were utilized to measure the wave-induced pore-pressure in
 119 the soil, as detailed in Fig. 1. Two wave height gauges (WHGs; model number: LYL-2)
 120 developed by Dalian University of Technology were located just above the PPTs. Far-field
 121 wave height was measured with the other two WHGs to guarantee the accuracy and reliability
 122 of the measured wave height and calculated wave length. The measurement accuracy of
 123 WHGs is 1 millimeter. The signals of WHGs and PPTs were multichannel synchronous
 124 sampled via the data acquisition card (NI USB-6211) with a sampling frequency of 100 Hz.
 125 An ADCP (model number: Vectrino-II; sampling volume: 0.085 cm^3 ; sampling frequency:
 126 100 Hz) was mounted to measure the flow velocity at the level of $0.5h$ (i.e., 0.25 m) above the
 127 sandy seabed near the PPTs.

128 Table 1. Index properties of test sands.

Mean size of sand grains	Geometric standard deviation	Coefficient of permeability	Void ratio	Relative density	Buoyant unit weight of soil
d_{50} (mm)	σ_g	k_s (m/s)	e	D_r	γ' (kN/m^3)

0.38	1.28	1.88×10^{-4}	0.771	0.352	9.32
------	------	-----------------------	-------	-------	------

129

130 **2.2. Test procedure and test conditions**

131

132

In general, the testing procedure was adopted as follows:

133

(1) The flume including the soil box was firstly emptied and cleaned.

134

(2) The PPTs were deaired and then saturated to ensure their argil-covers being free of air.

135

They were then installed at the specific locations with the support of a rack (see Fig. 1).

136

(3) The soil box was filled with clean water to a certain depth. The sand bed was carefully

137

prepared by means of sand-raining technique. The surface of the sand bed was leveled off

138

smoothly with a scraper.

139

(4) The flume was then filled slowly with water to a given depth (0.5 m in the present tests).

140

(5) For the tests of wave-only, the piston-type wave maker was activated and progressive

141

waves propagated from inlet onto the sandy seabed. For the tests of combined waves and

142

current, the current generator was firstly switched on and the flow velocity was gradually

143

increased to approach the target value. Thereafter, the wave maker was activated.

144

(6) The multichannel synchronous sampling system was then started to measure the multi-

145

physics parameters including wave height, pore-pressure and flow velocity.

146

Test conditions for investigating the wave-current induced excess pore-pressure in a

147

sandy seabed are summarized in Table 2. The mean water depth (h) was kept constant at 0.5

148

m. The wave period (T), wave height (H_0) and current velocity (U_c) were kept unchanged

149

during the test for the same run number. Note that, H_0 is the wave height under wave-only;

150

and U_c is the average velocity of the current without waves at the level of 0.25 m above the

151

sandy seabed. g is the gravitational acceleration. $\xi_0 (=H_0/L_0)$ is the wave steepness under

152

wave-only, where L_0 is the wave length under wave-only. The value of L_0 can be obtained

153

from the dispersion relationship:

154

$$L_0 = \frac{gT^2}{2\pi} \tanh(k_0 h) \quad (1)$$

155 where $k_0=2\pi/L_0$ is the wave number under wave-only. $\xi(=H/L)$ is the wave steepness
 156 considering the effect of current, where H and L are the wave height and wave length
 157 considering the effect of current, respectively. The values of H and L for calculating ξ are
 158 obtained from a theoretical expression for the variation of wave height and wave length with
 159 current velocity based on the linear theory of wave-current interaction (see Zou, 2004)

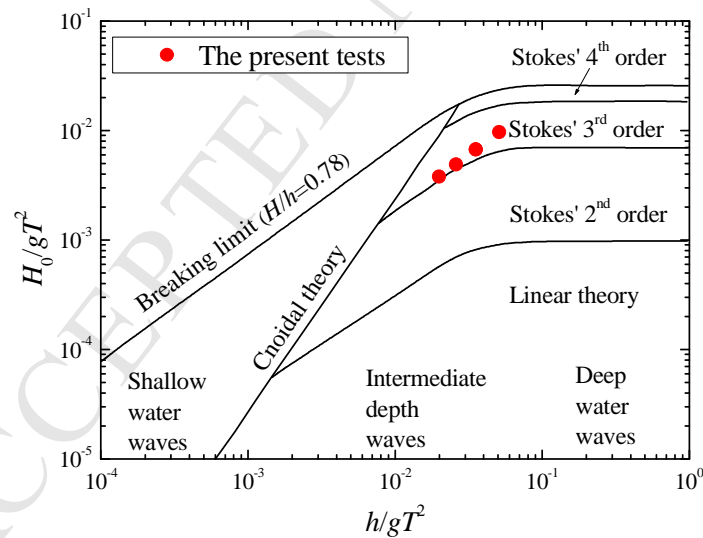
$$160 \quad \frac{H}{H_0} = 2(\chi + \chi^{1/2})^{-1/2} (1 + \chi^{1/2})^{-1/2} \quad (2a)$$

$$161 \quad \frac{L}{L_0} = \frac{1}{4}(1 + \chi^{1/2})^2 \quad (2b)$$

162 in which $\chi=1+4U_c/c_0$, $c_0=L_0/T$ is the wave velocity under wave-only.

163 According to the diagram of “the range of suitability of various wave theories” proposed
 164 by Lé Mehauté (1976), wave conditions of this study mainly fall in Stokes third-order wave
 165 theory zones (as shown in Fig. 2).

166



167

168

Fig. 2. Range of suitability of various wave theories (Lé Mehauté , 1976).

169

170 Supplementary tests were conducted to investigate the variation of wave height and
 171 wave length with the current velocity, and the profiles of flow velocity under various

172 hydrodynamic loading conditions. The test conditions for these supplementary tests are not
 173 elaborated in Table 2.

174

175 Table 2. Summary of test conditions for wave-current induced excess pore-pressure in a sandy seabed.

Run number	H_0 (cm)	T (s)	U_c (m/s)	H_0/gT^2	h/gT^2	ξ_0	ξ
1	9.5	1.2	0	0.0067	0.0354	0.046	0.046
2	9.5	1.2	0.1	0.0067	0.0354	0.046	0.037
3	9.5	1.2	0.2	0.0067	0.0354	0.046	0.031
4	9.5	1.2	0.3	0.0067	0.0354	0.046	0.027
5	9.5	1.2	-0.1	0.0067	0.0354	0.046	0.060
6	9.5	1.2	-0.2	0.0067	0.0354	0.046	0.084
7	9.5	1.2	-0.3	0.0067	0.0354	0.046	0.137
8	9.5	1.0	0	0.0097	0.0510	0.063	0.063
9	9.5	1.0	0.25	0.0097	0.0510	0.063	0.037
10	9.5	1.0	-0.25	0.0097	0.0510	0.063	0.167
11	9.5	1.2	0.25	0.0067	0.0354	0.046	0.029
12	9.5	1.2	-0.25	0.0067	0.0354	0.046	0.104
13	9.5	1.4	0	0.0049	0.0260	0.037	0.037
14	9.5	1.4	0.25	0.0049	0.0260	0.037	0.024
15	9.5	1.4	-0.25	0.0049	0.0260	0.037	0.077
16	9.5	1.6	0	0.0038	0.0199	0.031	0.031
17	9.5	1.6	0.25	0.0038	0.0199	0.031	0.020
18	9.5	1.6	-0.25	0.0038	0.0199	0.031	0.061

176

177 3. Results and discussions: Effects of imposing a current on waves

178 3.1. Variations of wave height and wave length

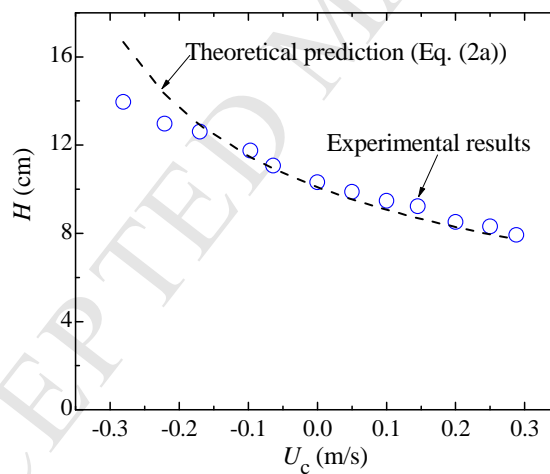
179 While waves and current coexist, the presence of a current will change the original wave
 180 height and wave length due to the interactions between waves and current. Fig. 3 shows the
 181 variation of the measured wave height and wave length with the velocity of the current (U_c ,
 182 refer to Fig. 4). The theoretical results calculated with Eq. (2) are also given in Fig. 3.

183 Theoretical solutions of Eq. (2) assume a uniform current and a deep water condition
 184 (i.e. $\tanh(k_0h) \rightarrow 1.0$), while in the experiments there exists an obvious boundary layer in the
 185 current (see Fig. 4) and the value of $\tanh(k_0h)$ is approximately 0.92. In spite of these two
 186 deviations from the actual experimental condition, Fig. 3 shows that the theoretical results are
 187 generally consistent with the experimental results in the current velocity range of $U_c > -0.1$ m/s.
 188 The wave height decreases and the wave length elongates significantly with increasing

189 velocity of a following-current. In contrast, the wave height is augmented and the wave length
 190 gets shorter with increasing velocity of an opposing-current, i.e., an opposing-current induces
 191 wave steepening.

192 Note that the theoretical solutions tend to overestimate the opposing-current-induced
 193 increase of wave height and decrease of wave length, while the velocity of the opposing-
 194 current is relatively large (e.g. $U_c = -0.22$ & -0.28 m/s). This might be due to the nonlinearity of
 195 wave-current interactions. Typical snapshots of the wave profiles under the conditions of $U_c =$
 196 0.20 m/s and $U_c = -0.30$ m/s are shown in Fig. 5(b) and 5(c), respectively. As a reference, a
 197 snapshot of the wave profile under wave-only is also given in Fig. 5(a). It is observed that the
 198 wave profiles under $U_c = -0.20$ m/s and $U_c = -0.30$ m/s are no longer sinusoidal. The surface of
 199 the waves are wrinkled up and apt to break, which implies a significant non-linearity of wave-
 200 current interactions (see Moreira and Peregrine, 2012).

201

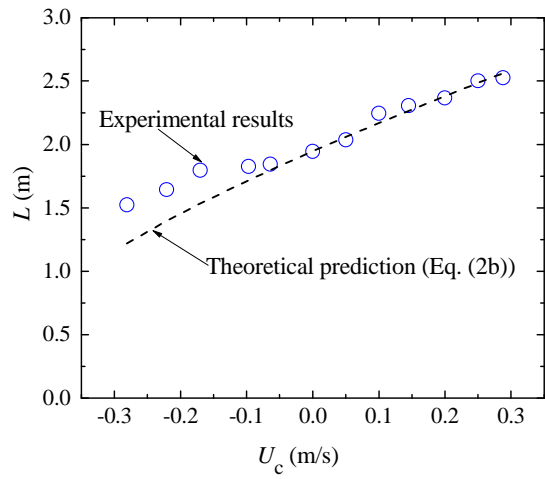


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203

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(a)

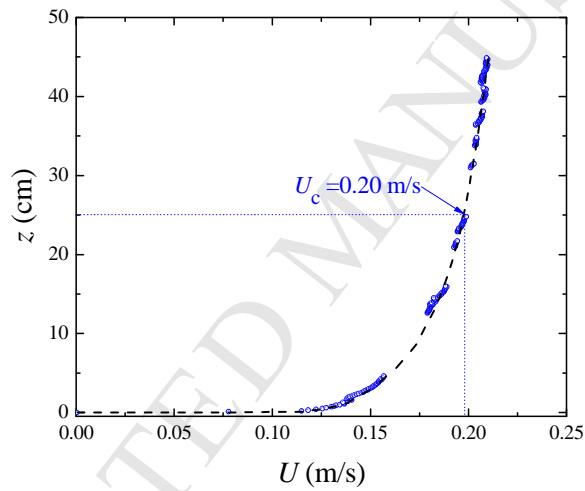


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206

207 Fig. 3. Effect of a current on the (a) wave height H ; and (b) wave length L . (Waves: $h=0.5$ m, $T=1.2$ s,

208 $H_0=10.2$ cm).

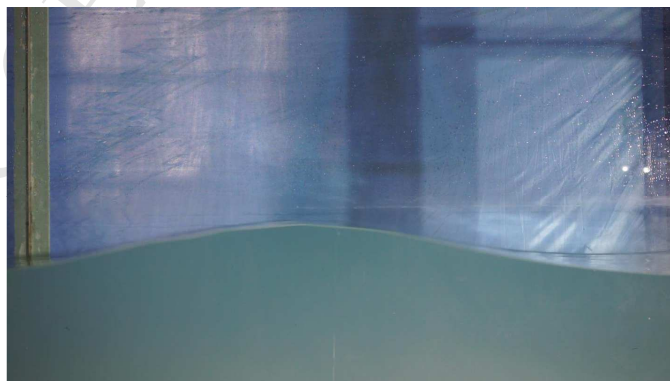
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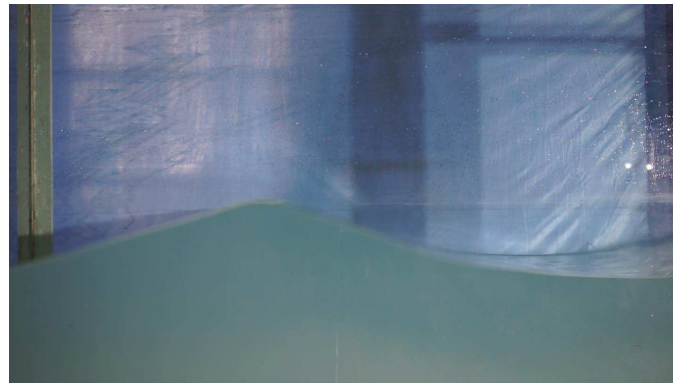
Fig. 4. Measured velocity profile of a unidirectional current.

212



(a) Wave-only

213
214

(b) Wave with an opposing current ($U_c = -0.20$ m/s)(c) Wave with an opposing current ($U_c = -0.30$ m/s)215
216217
218

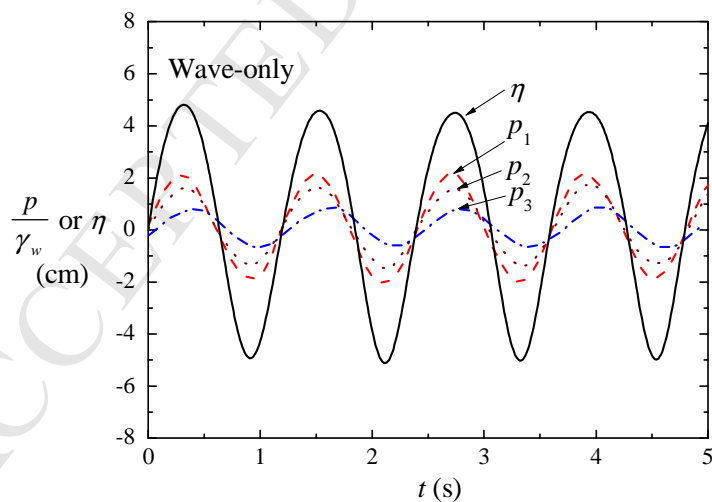
219 Fig. 5. Typical snapshots of the wave profiles under the conditions of: (a) wave-only; (b) $U_c = -0.20$ m/s;
220 and (c) $U_c = -0.30$ m/s. (Waves: $h = 0.5$ m, $T = 1.2$ s, $H_0 = 10.2$ cm)

221

222 3.2. Wave-induced pore-pressures

223 Figs. 6 (a) and 6(b) give the time series of the measured free surface elevation relative to
224 the static water level and corresponding excess pore-pressure responses at the same
225 measuring section under wave-only and waves with a following-current, respectively. As
226 shown in the figures, both the wave profile and wave-induced instantaneous pore-pressure
227 present a sinusoidal variation. No excess pore-pressure accumulation can be found in the
228 present sandy seabed under the examined hydrodynamic loads. This absence of pore-pressure
229 accumulation should be attributed to the relatively large permeability of the soil ($k_s = 1.84 \times 10^{-4}$
230 m/s, $c_v = 0.66$) and apparently smaller wave-induced shear stress in the soil compared with
231 that in a typical prototype condition (Jeng and Seymour, 2007 or Figure 2.11 in Jeng (2018)).

232 The wave-induced pore-pressure at the surface of the sandy seabed (p_1) has the same
 233 phase with the free surface elevation. An evident phase lag can be observed among the pore-
 234 pressure responses measured at three different soil depths (p_1 , p_2 and p_3 , refer to Fig. 1 for the
 235 detailed locations). The analysis of Yamamoto et al. (1978) indicates that no phase lag would
 236 occur in a completely saturated infinite seabed, because the wave-induced pore pressures and
 237 effective stresses are independent of soil characteristics in such a condition. However, this
 238 conclusion was based on the case of infinite seabed. As reported in Jeng and Hsu (1996), the
 239 conclusion from Yamamoto et al. (1978) is no longer valid for a seabed finite thickness,
 240 because the soil characteristics directly affect the pore pressures and effective stresses and
 241 cause minor phase lag even for nearly saturated seabed. This physical process is attributed to
 242 the multi-phase flow in a porous medium. Furthermore, this phenomenon only occurs in fine
 243 sand such as the present tests (Jeng and Hsu, 1996). The comparison between Fig. 6(a) and
 244 6(b) indicates that superimposing a following-current upon waves has a minor effect on the
 245 phase lag of the excess pore-pressure responses in the sandy seabed.
 246

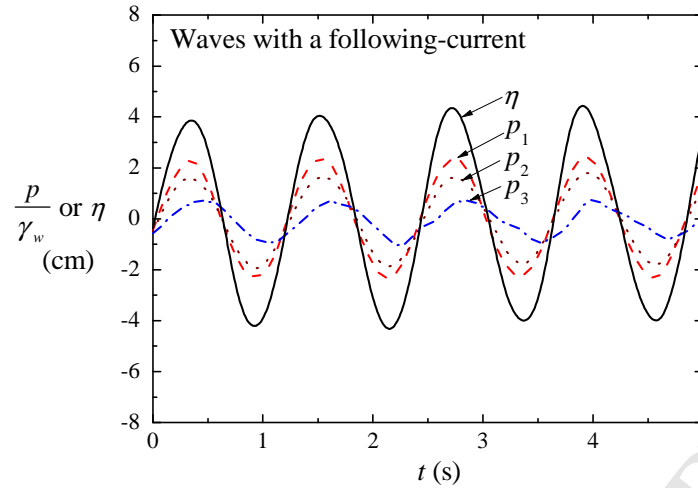


(a)

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248

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(b)

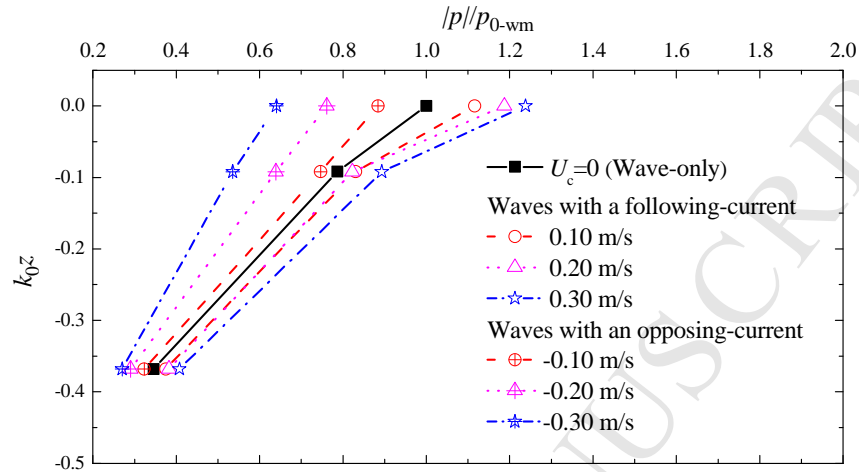
Fig. 6. Time series of free surface elevation relative to the static water level (η) measured with WHG-III and excess pore-pressure measured with PPT1 (p_1), PPT2 (p_2) and PPT3 (p_3): (a) wave-only; and (b) waves with a following-current. (Waves: $h=0.5$ m, $T=1.2$ s, $H_0=9.5$ cm; current: $U_c=0.25$ m/s)

Fig. 7 illustrates the distributions of the excess pore-pressure amplitude ($|p|/p_{0-wm}$, where p_{0-wm} is the measured amplitude of the wave-induced pore-pressure at mudline under wave-only) along the soil depth (k_0z) under different combinations of waves and current loadings. It is shown that if a following-current is superimposed onto waves, the excess pore-pressure amplitudes in the sandy seabed generally increase. In contrast, an opposing-current would decrease the excess pore-pressure amplitudes in the sandy seabed. Specifically, the increment/reduction of the pore-pressure amplitude at mudline due to a following/opposing-current of $|U_c|=0.30$ m/s can be up to 35%/24%.

It is also observed from Figure 7 that the excess pore-pressure gradients in the seabed would generally be increased/reduced by a following/opposing-current. These results indicate that the liquefaction or partial liquefaction is more likely to occur under combined waves and following-current loading, while the opposing-current is beneficial to prevent the seabed to liquefying. That is, a following-current might be a potential risk for the safety of offshore structures. Moreover, the excess pore-pressure gradients would exert lifting force onto the sand grains under the wave-troughs and thereby might bring the sand more susceptible to scour. Although the value of the gradient variation caused by superimposing a current are not

272 large in the present experiments, it could become significant in a real ocean environment
 273 where the wave height and wave period can be more than 10 times greater than those in the
 274 flume experiment.

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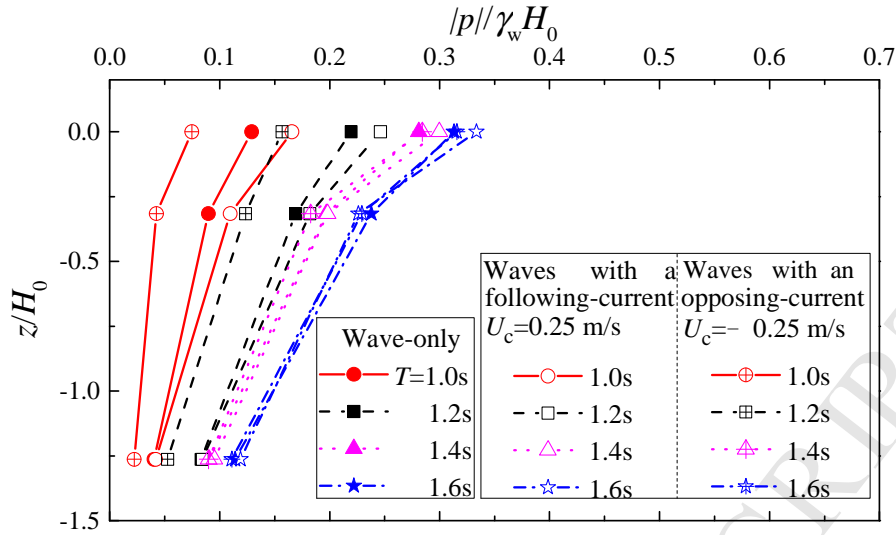
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278 Fig. 7. Effects of a following/opposing-current with various velocities on the vertical distributions of
 279 the excess pore-pressure amplitude along the soil depth. ($T=1.2$ s, $H_0=9.5$ cm)

280

281 The effects of a following/opposing-current ($|U_c|=0.25$ m/s) on the distributions of the
 282 excess pore-pressure amplitude under various wave periods are compared in Fig. 8. Note that
 283 unlike Fig. 7, the normalized excess pore-pressure amplitude is expressed as $p/\gamma_w H_0$ and the
 284 normalized soil depth is expressed as z/H_0 in Fig. 8, since the values of p_{0-wm} and k_0 vary with
 285 wave period. It is indicated that as wave period increases, the current-induced difference of
 286 the excess pore-pressure amplitude gradually becomes small. Under the conditions of $T=1.0$ s
 287 and 1.2 s, the magnitude of the following-current induced enlargement of the excess pore-
 288 pressure amplitude is obviously smaller than the opposing-current induced reduction of the
 289 excess pore-pressure amplitude. Under the conditions of $T=1.4$ s and 1.6 s, the effects of the
 290 following-current are small but observable, while the effects of the opposing-current seem to
 291 be negligible.

292



293

294 Fig. 8. Comparison of vertical distributions of the excess pore-pressure amplitude along soil depth
 295 between the conditions of wave-only, waves with a following-current ($U_c=0.25$ m/s), and waves with
 296 an opposing-current ($U_c=-0.25$ m/s) for various wave periods. ($H_0=9.5$ cm)

297

298 3.3. Comparisons between experimental and analytical results

299 As aforementioned, Zhang et al. (2013) proposed an analytical solution for the wave-
 300 induced pore-pressure responses in the seabed under combined waves and current. In their
 301 model, the third-order approximation of the wave-current interactions proposed by Hsu et al.
 302 (2009) was employed for the dynamic wave pressure acting on the seabed. As shown in Fig.2,
 303 the third-order Stokes wave theory is in accord with the practical wave conditions in the
 304 present flume tests. The dynamic wave pressure acting on the seabed is expressed as (Ye and
 305 Jeng, 2012)

$$306 \quad P_b(x,t) = P_1 \cos(k_0 x - \omega t) + P_2 \cos 2(k_0 x - \omega t) + P_3 \cos 3(k_0 x - \omega t) \quad (3)$$

$$307 \quad \text{where } P_1 = \frac{\rho_f g H_0}{2 \cosh k_0 h} \left\{ 1 - \frac{\omega_0 k_0^2 H_0^2}{2(U_c k_0 - \omega_0)} \right\}, \quad P_2 = \frac{3\rho_f H_0^2}{8} \left\{ \frac{\omega_0(\omega_0 - U_c k_0)}{2 \sinh^4(k_0 h)} - \frac{gk_0}{3 \sinh 2k_0 h} \right\},$$

$$308 \quad P_3 = \frac{3\rho_f k H_0^3 \omega_0(\omega_0 - U_c k_0)}{512} \frac{9 - 4 \sinh^2(k_0 h)}{\sinh^7(k_0 h)}, \text{ and } \rho_f \text{ is the water density. The dispersion}$$

309 relationship is given as

$$310 \quad \omega = \omega_0 + (k_0 H_0)^2 \omega_2 \quad (4)$$

$$311 \quad \text{where } \omega_0 = U_c k_0 + \sqrt{g k_0 \tanh k_0 h} \quad \text{and} \quad \omega_2 = \frac{9 + 8 \sinh^2(k_0 h) + 8 \sinh^4(k_0 h)}{64 \sinh^4(k_0 h)} (\omega_0 - U_c k_0).$$

312 Taking Eq. (3) as a boundary condition at the mudline and based on the quasi-static Biot's
 313 consolidation equations (Biot, 1941), the excess pore-pressure for a uniform and isotropic
 314 seabed can be derived as (Zhang et al., 2013)

$$315 \quad P = \sum_{m=1}^3 \frac{P_m}{1-2\mu} \left[(1-2\mu-\alpha) C_{1m} e^{mk_0 z} + \frac{\delta_m^2 - m^2 k_0^2}{mk_0} (1-\mu) C_{2m} e^{\delta_m z} \right] e^{im(k_0 x - \omega t)} \quad (5)$$

$$316 \quad \text{where } \alpha = \frac{(1-2\mu)n\beta}{n\beta + (1-2\mu)/G}, \quad \delta_m^2 = m^2 k_0^2 - \frac{im\omega\rho_f g}{k_s} \left\{ n\beta + \frac{1-2\mu}{2G(1-\mu)} \right\},$$

$$317 \quad C_{1m} = \frac{\delta_m - \delta_m \mu + mk_0 \mu}{\delta_m - \delta_m \mu + mk_0 \mu + mk_0 \alpha}, \quad C_{2m} = \frac{mk_0 \alpha}{(\delta_m - mk_0)(\delta_m - \delta_m \mu + mk_0 \mu + mk_0 \alpha)}, \quad \mu \text{ is the}$$

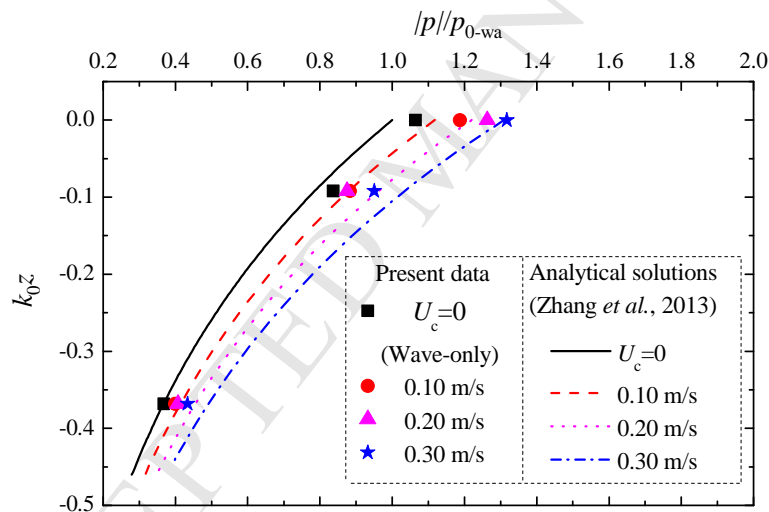
318 Poisson ratio, n is the soil porosity, β is the compressibility of pore fluid, G is the shear
 319 modulus of the soil, and k_s is the soil permeability.

320 Fig. 9 gives the comparison of the vertical distributions of the excess pore-pressure
 321 amplitude (p/p_{0-wa} , where p_{0-wa} is the analytically calculated amplitude of the wave-induced
 322 pressure fluctuation at the mudline under wave-only) along the soil depth ($k_0 z$) between
 323 experimental results and analytical solutions calculated with Eq. (5). The values of the input
 324 parameters for the analytical solutions are shown in Table 2. The degree of saturation is a key
 325 influencing factor for determining the distribution of the wave-induced pore-pressure (Okusa,
 326 1985; Sakai et al., 1992). Nevertheless, the specific value of the degree of saturation is
 327 difficult to measure accurately (Michallet et al., 2009). In the present comparison study, the
 328 experimental data for the conditions of wave-only are utilized to calibrate the value of the
 329 degree of saturation (see Figs. 9 and 11(a)). It is proved that the value of $S_r=0.995$ would
 330 generally make the analytical results coincide well with the experimental results. Note that the
 331 amplitude of the wave-induced pressure fluctuation at the mudline is only influenced by wave
 332 parameters and irrelevant to the seabed properties. As such, the deviations of the wave-

333 induced pressure fluctuation at the mudline between analytical and experimental results are
 334 intrinsic and unaffected by the calibrated value of the degree of saturation.

335 Fig. 9 shows that the variation trend of the distributions of the excess pore-pressure
 336 amplitude calculated with the analytical solution is generally consistent with the experimental
 337 data. By comparing Fig. 9(a) with Fig. 9(b), it can be seen that certain deviations exist
 338 between the flume results for excess pore-pressure and the analytical solutions, which is
 339 nonnegligible especially for the opposing-current case. The relatively larger deviations for the
 340 opposing-current cases can be mainly attributed to the intrinsic difference between the
 341 measured amplitude of the wave-induced pressure fluctuation at the mudline and the
 342 analytical one (see Fig. 9(b)).

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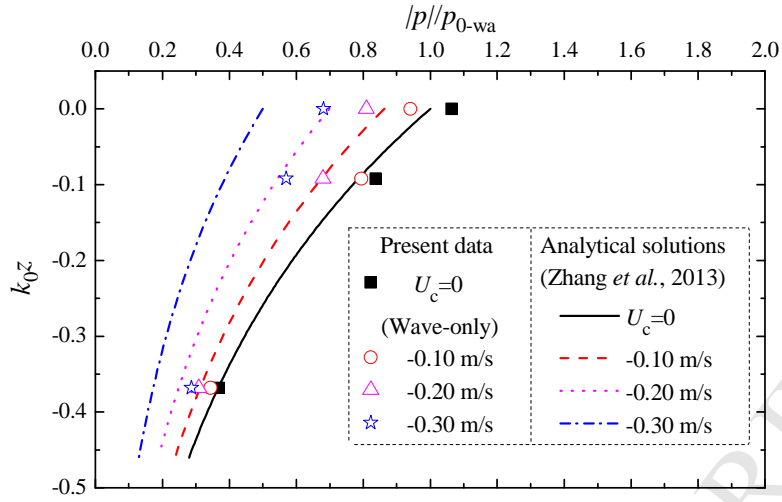


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345

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(a)



(b)

Fig. 9. Comparisons of the vertical distributions of the excess pore-pressure amplitude along the soil depth between experimental results and analytical solutions under (a) waves with a following-current (various current velocities); and (b) waves with an opposing-current (various current velocities). ($h=0.5$ m, $T=1.2$ s, $H_0=9.5$ cm).

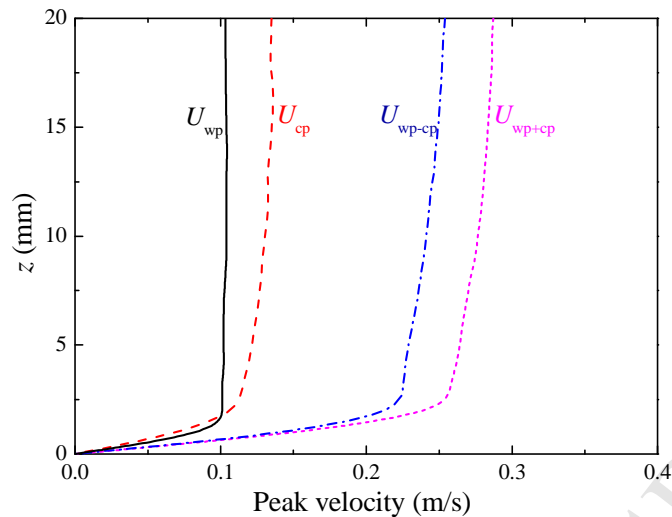
Table 2. Values of the input parameters for the analytical solutions.

Parameters		Values
Seabed properties	Coefficient of permeability k_s (m/s)	1.88×10^{-4}
	Shear modulus G (MPa)	10.0
	Void ratio e	0.771
	Poisson ratio of soil ν	0.30
	Degree of saturation S_r	0.995
Wave parameters	Water depth h (m)	0.5
	Wave height H_0 (cm)	9.5
	Wave period T (s)	1.2 (Various in Fig. 11)

To clarify the fact that the difference between the measured amplitude and the analytically calculated amplitude of the wave-induced pore-pressure at mudline under waves with an opposing-current is larger than that under waves with a following-current, supplementary measurements for the velocity profiles of the boundary layer under conditions

360 of wave-only, current-only, waves with a following-current, and waves with an opposing-
361 current, were conducted. The flow velocity measured at $z=0.25$ m under current-only (U_c) is
362 0.20m/s. The wave period is 1.5 s and the wave height under wave-only (H_0) is 7.2cm. The
363 sampling duration of each case was 5 min (approximately 200 wave cycles) with a sampling
364 frequency of 100 Hz in order to have a statistically time independent average velocity (or
365 phase-averaged velocity). The measured (phase-averaged) peak velocity profiles along the
366 water depth under wave-only (U_{wp}), current-only (U_{cp}), waves with a following-current
367 (U_{wp+cp}), and waves with an opposing-current (U_{wp-cp}), are shown in Fig. 10. The actual
368 velocity profiles under waves and a following-current were found obviously different from
369 those suggested by a linear superposition of wave-only and current-only velocities.

370 While analytically calculating the amplitude of the wave-induced pressure fluctuation at
371 the mudline after Hsu et al. (2009) and Ye and Jeng (2012), the current was assumed to be
372 uniform and the measured flow velocity at $z=0.25$ m under current-only (U_c) was chosen as the
373 input value of the current velocity. However, there exists a significantly thick boundary layer
374 in the current (see Fig. 4) and thus the input velocity at $z=0.25$ m for the analytical calculation
375 would be much greater than the actual average current velocity. This could be the reason that
376 the analytically calculated amplitude of the wave-induced pressure fluctuation at the mudline
377 is much smaller than the measured one under waves with an opposing-current. For the cases
378 of waves with a following-current, Fig. 10 indicates that the maximum flow velocities are
379 much larger than those for opposing-current cases and also the sum of wave and current,
380 which is consistent with the result of Kemp and Simons (1982; 1983), Olabarrieta et al.
381 (2010) and Qi and Gao (2014). This enhancement of flow velocities near the seabed could
382 somehow compensate the difference between the input current velocity for analytical
383 calculation and the actual current velocity considering boundary layer. As a result, the
384 deviation of the analytically calculated amplitude of the wave-induced pore-pressure at
385 mudline from the measured one under waves with a following-current is much smaller
386 compared with the cases of waves with an opposing-current.



387

388 Fig. 10. Comparisons of the profiles for peak velocities between the conditions of wave-only,
 389 only, waves with an opposing-current and waves with a following current. (Waves: $h=0.5$ m, $T=1.2$ s,
 390 $H_0=10.2$ cm; current: $U_c=0.20$ m/s)

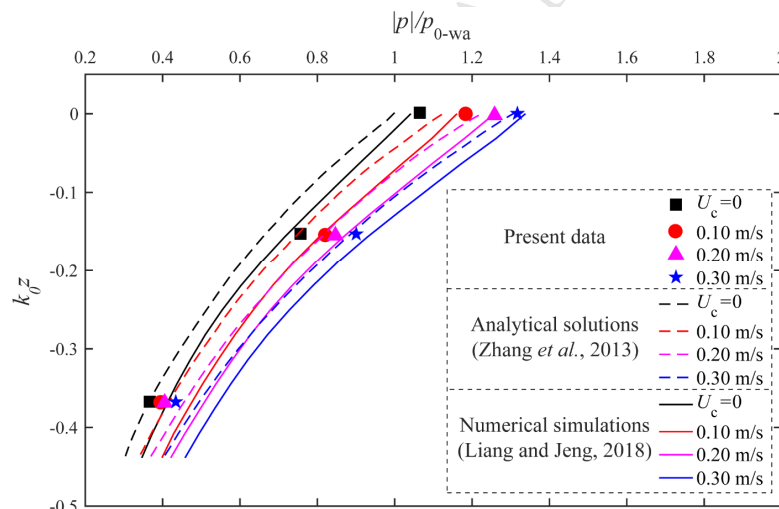
391

392 The third-order approximation proposed by Hsu et al. (2009) was based on the potential
 393 flow theory, which cannot describe the boundary layer flow along the seabed surface. To
 394 further confirm the aforementioned conjectured effects of the boundary layer on the combined
 395 wave-current induced pore-pressure responses in the seabed, several numerical simulations in
 396 which the boundary layer of the wave and/or current can be generated were carried out. In the
 397 numerical model, a hydrodynamic model based on the finite volume method and volume-
 398 averaged Reynolds-averaged Navier-Stokes equation (del Jesus et al., 2012) is developed to
 399 investigate the interactions between the third-order wave and shear current, while the quasi-
 400 static Biot equation (Biot, 1941) is adopted to describe the mechanical behaviour of a
 401 hydraulically isotropic porous elastic seabed in the open-source CFD toolbox Open-FOAM.
 402 The dynamic wave pressure extracted from the hydrodynamic model is utilized as the
 403 boundary condition at the seabed surface for the seabed model. In this model, the bottom
 404 boundary layer near the seabed surface is included in the numerical simulation. More detailed
 405 descriptions with respect to the numerical model can be found in Liang and Jeng (2018).

406 The comparison of the vertical distributions of the excess pore-pressure amplitude along
 407 the soil depth between experimental results, analytical solutions calculated with Eq. (5) and

408 aforementioned numerical simulations are provided in Fig. 11. The values of the input
 409 parameters for the analytical solutions and numerical simulations are shown in Table 2. The
 410 most important finding from Fig. 11 is that the magnitude of the pore-pressure amplitude at
 411 the seabed obtained by numerical simulations are generally much closer to the experimental
 412 data compared with the results of analytical solutions, especially under the conditions of
 413 waves with an opposing-current. This observation explicitly highlights the effects of
 414 boundary layer on the combined wave-current induced pore-pressure responses in the seabed.
 415 It is noted that observable deviation of the pore-pressure amplitude at the seabed still exists
 416 between the experimental results and the numerical simulations, mostly due to the difficulties
 417 of reproducing all the important characteristics of wave and/or current boundary layer and
 418 wave-current interactions in the present numerical hydrodynamic model.

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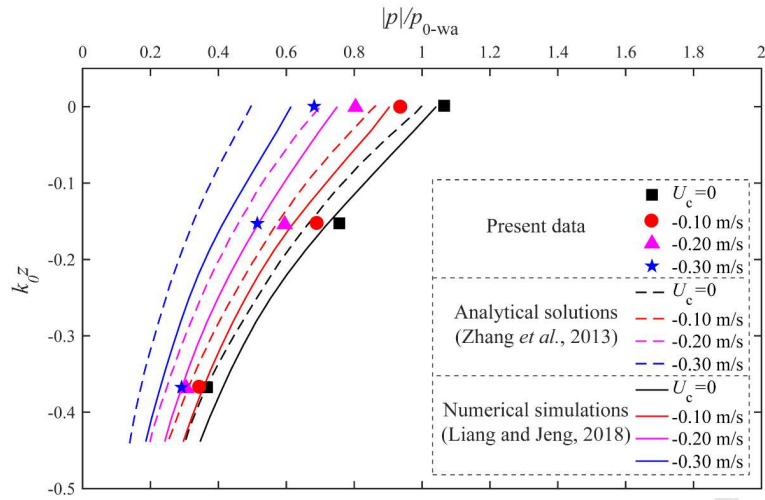


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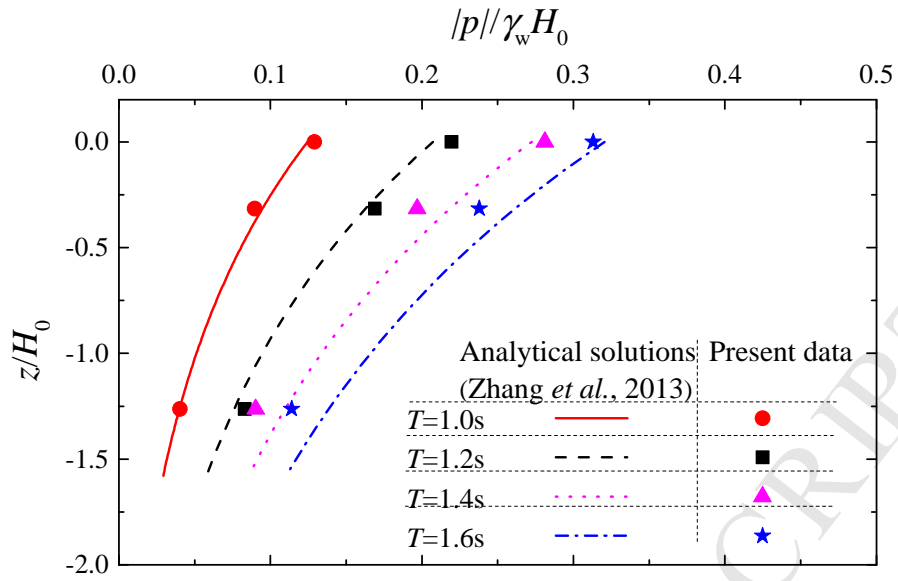
(a)



(b)

Fig. 11. Comparisons of the vertical distributions of the excess pore-pressure amplitude along the soil depth between experimental results, analytical solutions and numerical simulations under (a) waves with a following-current (various current velocities); and (b) waves with an opposing-current (various current velocities). ($h=0.5$ m, $T=1.2$ s, $H_0=9.5$ cm).

The comparisons between the measured vertical distributions of the excess pore-pressure amplitude and the analytical ones along the soil depth under various wave periods are given for the cases of wave-only, waves with a following-current and waves with an opposing-current in Figs. 12(a), 12(b) and 12(c), respectively. It is shown that in terms of the amplitude of the wave-induced pressure fluctuation at the mudline, the analytical and experimental results are generally consistent under wave-only (see Fig. 12(a)). Nevertheless, the analytical ones are a bit larger than the experimental ones under waves with a following current, whilst much smaller than the experimental ones under waves with an opposing-current. Focusing on the attenuation rate of the pore-pressure along the soil depth (or the general profile of the pore-pressure normalized with pore-pressure at the seabed surface), Figs. 12(a) and 12(b) indicate that the analytical and experimental results generally match well for the cases of wave only and waves with a following-currents.

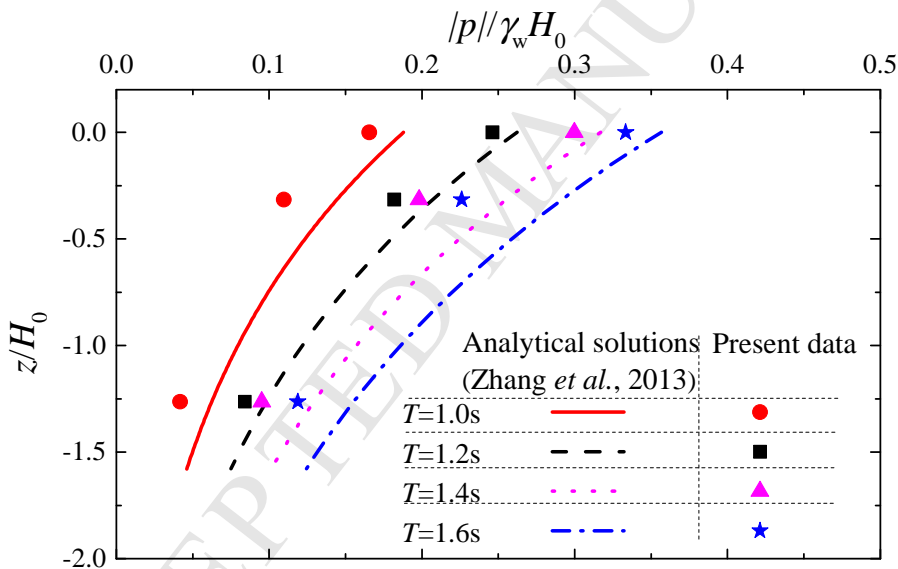


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(a)

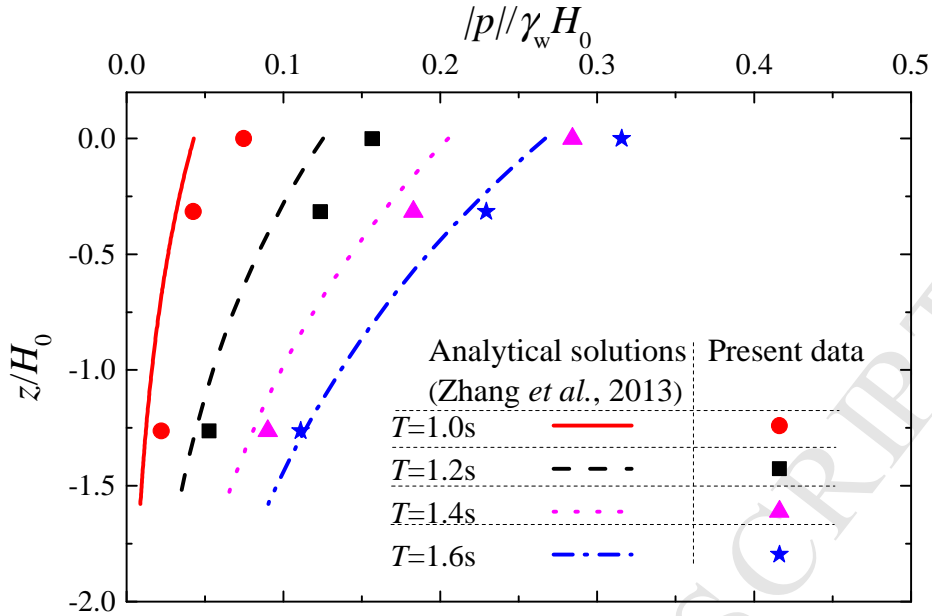


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(b)

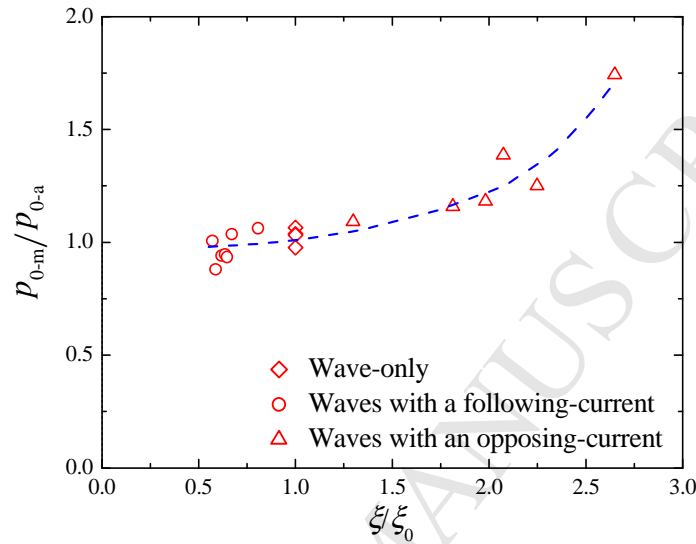


(c)

449
 450
 451
 452 Fig. 12. Comparisons of the vertical distributions of the maximum pore-pressure along the soil depth
 453 between experimental results and analytical solutions under (a) wave-only; (b) waves with a following-
 454 current ($U_c=0.25$ m/s); and (c) waves with an opposing-current ($U_c=-0.25$ m/s). (Various wave periods,
 455 wave height $H_0=9.5$ cm)

456
 457 The present flume observations indicated that superimposing a steady current on waves
 458 could significantly alter the pressure on the sand-bed surface. The deviation of the measured
 459 excess pore-pressure from analytical solutions can be mainly attributed to the difference
 460 between the measured pore pressure amplitude at the mudline (denoted as p_{0-m}) and the
 461 analytical one (p_{0-a}). Fig. 13 gives the variation of p_{0-m}/p_{0-a} with the ratio of wave
 462 steepness under combined waves and current to that under wave-only (ξ/ξ_0). ξ/ξ_0
 463 indicates the effect of a current on the wave steepness. A general and consistent trend, i.e., the
 464 value of p_{0-m}/p_{0-a} is obviously larger than 1.0 and keep increasing with increasing value of
 465 ξ/ξ_0 when $\xi/\xi_0 > 1$ (i.e. opposing-current cases) while the value of p_{0-m}/p_{0-a} is
 466 generally around 1.0 when $\xi/\xi_0 \leq 1$ (i.e. wave-only and following-current cases), is
 467 indicated in Fig. 13. This trend implies that in terms of the combined wave-current induced

468 excess pore-pressure in a sandy seabed, a favorable prediction by analytical solution should
 469 be expected for following-current cases and smaller pore-pressure amplitudes would be
 470 obtained for opposing-current cases, which has been confirmed by Figs. 9 and 12. While
 471 evaluating the potentially enhanced risk for the safety of offshore structures by a following-
 472 current, the analytical solution would provide a conservative/safe prediction.



473

474 Fig. 13. The variation of p_{0-m}/p_{0-a} with ξ/ξ_0 ($H_0/gT^2 = 0.0038 \sim 0.0097$,

475

$$h/gT^2 = 0.0199 \sim 0.0510).$$

476 4. Conclusions

477 The co-existence of waves and current in offshore environments is a common scenario in
 478 coastal zones. A series of flume tests were conducted to investigate the effect of imposing
 479 following/opposing current upon waves on the excess pore-pressure in the sandy seabed. This
 480 study provide the first set of comprehensive experimental data for the pore-water pressures in
 481 a porous seabed due to combined waves and currents. Based on flume observations and
 482 comparisons with the existing theoretical solution, the following conclusions are drawn:

- 483 (1) The essential difference of the pore-pressure responses between wave-only condition and
 484 combined wave-current condition is due to the different boundary conditions of pressure
 485 at the seabed surface. The excess pore-pressure amplitudes are increased for the
 486 following-current case, but reduced for the opposing-current case. For the examined

487 value range, such wave-current combination effect becomes more significant for shorter
488 wave periods. The excess pore-pressure gradients in the seabed would be
489 enhanced/reduced by a following/opposing-current.

490 (2) The variation trend of the excess pore-pressure distribution in the present flume
491 observations is consistent with that of the existing theoretical solutions. Nevertheless,
492 certain deviations exist between the flume results for excess pore-pressure and the
493 analytical solutions, which is nonnegligible especially for the opposing-current case. The
494 apparently larger deviations under waves with an opposing-current can be mainly
495 attributed to the intrinsic difference between the analytically calculated amplitude of the
496 wave-induced pressure fluctuation at the mudline and the measured one.

497 (3) Measurements for the velocity profiles of the boundary layer shows that the maximum
498 flow velocities under waves with a following-current are much larger than those for
499 opposing-current cases and also the sum of wave and current. This enhancement of flow
500 velocities near the seabed under waves with a following-current could compensate the
501 overestimated input current velocity for analytical calculation induced by a boundary
502 layer. Therefore, the deviation between the analytically calculated amplitude of the wave-
503 induced pressure fluctuation at the mudline and the measured one under waves with a
504 following-current is much smaller than the cases of waves with an opposing-current. The
505 effects of boundary layer on the combined wave-current induced pore-pressures in the
506 seabed are highlighted by supplementary numerical simulations.

507 (4) A general and consistent variation trend of p_{0-m}/p_{0-a} with the ratio of wave steepness
508 under combined waves and current to that under wave-only is indicated based on the
509 present experimental results. A favorable prediction by the analytical solution would be
510 expected for following-current cases and smaller pore-pressure amplitudes would be
511 obtained for opposing-current cases. Therefore, while evaluating the potentially
512 enhanced risk for the safety of offshore structures by a following-current, the analytical
513 solution would provide a conservative/safe prediction.

514

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522

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Highlights:

- ✧ Flume observations of combined wave-current induced excess pore-pressure in a sand-seabed;
- ✧ Comparisons of wave-current induced pore-pressure between flume results and analytical solutions;
- ✧ Effects of non-linearity of wave-current interactions on the excess pore-pressure responses;
- ✧ Deviation between the flume results for excess pore-pressure and the analytical solutions has been identified.